Application of continuous FerryBox measurements to oxygen and carbon fluxes in the North Sea

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Motivation

- CO₂ in the atmosphere → climate change
- to which extent it is dissolved in the ocean
 - \rightarrow primary production
 - \rightarrow acidification
- continuous O₂ measurements in FerryBox system
- O₂ serves as a proxy for primary production
 - quantification of carbon fluxes at air-sea interface





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> Air-sea flux of oxygen expressed by gas concentration anomaly and gas exchange velocity:

$$F = k_w \cdot \Delta O_2$$

- Gas concentration anomaly is difference of observed concentration and saturation concentration
- $\succ k_w$ is parameterized by wind speed (in 10 meter height) and the dimensionless Schmidt number Sc.

$$Sc = \frac{\mu}{D} = A - bt + Ct^2 - Dt^3$$

 μ is kinematic viscocity of water, D is diffusion coeffcient of the gas

general parameterization term:

$$k_w = a \cdot Sc^n \cdot U^b$$

➢ f.e. Wanninkhof, 1992:

$$k_w = 0.31 \cdot (S^{C}/_{660})^{-1/2} \cdot U^2$$

Defining carbon flux rates via Redfield molar ratio of 0.77 (C:O)

Data



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- FerryBox data of Tor Dania 04/2011-04/2012
 - Dissolved oxygen
 - > Water temperature
 - > Salinity
- ECMW ERA-Interim reanalysis data, 0.75° grid, 6-hourly
- BSH cmod model 0.02° grid, 1 hour time resolution





Ferrybox measurements of dissolved oxygen



Part of BSHcmod grid

FerryBox measurements Tor Dania 2011-2012





Dissolved oxygen



Water temperature

FerryBox measurements & model output







Results





Parameter tests



- Impact of choice of parameterization scheme
- Impact of high wind speeds, exclusion of winds > 12 m/s
- Wind speed: ERA-Interim vs BSH cmod model

Parameterization solutions



Parameterisation schemes for exchange velocity of air-sea flux of oxygen in dependence of wind speed U_{10} :

W-92: Wanninkhof, 1992:

$$k_w = 0.31 \cdot (S_c/_{660})^{-1/2} \cdot U^2$$

WM-99: Wanninkhof & McGillis, 1999: $k_w = 0.0283 \cdot u^3 \cdot (\frac{Sc}{660})^{-1/2}$ N-00: Nightingale, 2000: $k_w = (0.2222u^2 + 0.333u)(\frac{Sc}{660})^{-1/2}$ LM-86: Liss & Merlivat, 1986: $k_w = 0.17u \cdot (\frac{Sc}{660})^{-2/3}$ $k_w = (2.85u - 9.65) \cdot (\frac{Sc}{660})^{-1/2}$ $k_w = (5.9u - 49.3) \cdot (\frac{Sc}{660})^{-1/2}$ W-05: Woolf, 2005:

$$k_w = (56.52\sqrt{C_d}u + 2.5 \cdot 10^{-4}u^{4.04}) \cdot (\frac{Sc}{660})^{-1/2}$$

Parameterizations



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Wind forcing by ERA-Interim

Exclusion of high winds





Wind forcing by ERA-I, Wanninkhof (1992) parameterization

Wind input BSHcmod vs ERA-I

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Carbon fluxes: Wind input BSH cmod vs ERA-I





- Results of flux calculations for one annual cycle from Tor Dania dataset
- ➢ In general, southern North Sea is sink for carbon during growing season
- > Carbon fluxes reach up to 600 gC/m² in one year, located in the German Bight
- Tests show
 - variation for different parameterizations of exchange velocity
 - > High impact of high winds, although they occurr mainly in winter
 - Coarse ERA-I dataset affect lower carbon fluxes near the coast

- dataset over longer timescale
 - \blacktriangleright more than one seasonal cycle \rightarrow annual variability (Petersen et al., 2011)
- calculations for other ferrybox routes,
 - Iong timescales
 - continuous data sets
- Extent study of parameter reliability

Thank you!



- Transect data of Tor Dania: every 2-3 days a data set at one point of the transect
- Create grid of 0.05°E (x-axis) x 7 days (y-axis)
- Interpolate model wind data on that grid

Carbon fluxes depending on wind forcing





- Wanninkhof (1992) parameterization
- BSH cmod data available until end of 2011